R005-13 A 会場 :11/24 PM2 (15:30-18:15) 16:45~17:00

#西山 尚典 ^{1,2)}, 鍵谷 将人 ³⁾, Bag Tikemani¹⁾, 津田 卓雄 ⁴⁾, 岩佐 祐希 ⁵⁾, 小川 泰信 ^{1,2)}, Sigernes Fred⁶⁾ ⁽¹ 極地研, ⁽² 総研大, ⁽³ 東北大・理・惑星プラズマ大気研究センター, ⁽⁴ 電通大, ⁽⁵ 計量標準総合センター/産業技術総合研 究所, ⁽⁶University centre in Svalbard

Thermospheric orthohelium variations associated with a moderate storm on February 2023: the NIRAS-2 observations at Longyearbyen

#Takanori Nishiyama^{1,2)}, Masato Kagitani³⁾, Tikemani Bag¹⁾, Takuo Tsuda⁴⁾, Yuki Iwasa⁵⁾, Yasunobu Ogawa^{1,2)}, Fred Sigernes⁶⁾

⁽¹National Institute of Polar Research, ⁽²The Graduate University for Advanced Studies, SOKENDAI, ⁽³Planetary Plasma and Atmospheric Research Center, Graduate School of Science, Tohoku University, ⁽⁴University of Electro-Communications, ⁽⁵NMIJ/AIST, ⁽⁶University centre in Svalbard

Metastable orthohelium, $He(2^{3}S)$, is known to have airglow emissions by solar resonance scattering at wavelength of 388.9 nm (3³P-2³S) and 1083 nm (2³P-2³S). The emissions from the terrestrial atmosphere have been first reported more than 60 years ago during strong twilight and sunlit aurora over Moscow (Mironov et al., 1959; N. Shefov, 1961). Transition from ground state helium, $He(1^{1}S)$, to $He(2^{3}S)$ is mainly caused by photoelectron impacts with energy larger than 19.8 eV (de Heer & Jansen, 1977). $He(2^{3}S)$ layer is confined in the upper thermosphere and in the lower exosphere. Photoelectron impact occurs with a maximum production rate around 250 km. Generated $He(2^{3}S)$ diffuses upward, but rapidly disappears above 500 km due to photo ionisation. Below 250 km, quenching due to Penning ionization is significant. As a result, $He(2^{3}S)$ layer can be formed from 250 km to 800 km or above (Bishop & Link, 1993; Waldrop et al., 2005). That's why $He(2^{3}S)$ is an unique and attractive candidate for remote sensing target of the neutral atmospheric dynamics in this region.

This study presents a time variability of $He(2^{3}S)$ in polar region associated with a geomagnetic storm for the first time. Continuous 17-nights dataset of $He(2^{3}S)$ airglow brightness at 1083 nm was obtained from a short-wavelength infrared imaging spectrograph (NIRAS-2) installed at Longyearbyen, Svalbard (78.1° N, 16.0° E). The observed $He(2^{3}S)$ airglow brightness clearly displayed responses to a geomagnetic storm in different time scales. The $He(2^{3}S)$ airglow brightness has began to decrease sharply within an hour of sudden commencement of the storm, and it was gradually decreased over the next few days and then recovered slowly. It was nicely agreed to helium density variations at 500-km altitude calculated by MSIS. The depletion of $He(2^{3}S)$ was mainly caused by enhanced Penning ionization due to upwelling N₂ from the lower atmosphere; this was consistent with decreased O/N₂ ratio in MSIS and TIMED/GUVI measurements and electron density depletion in F region observed by EISCAT Svalbard Radar (ESR). Additionally, sudden increases in $He(2^{3}S)$ airglow brightness were clearly found associated with intermittent electron precipitations observed by the ESR. Therefore, direct impact by precipitating electron injected from the space to the polar upper atmosphere can play significant roles in production of $He(2^{3}S)$. The NIRAS-2 measurements have successfully demonstrated that column density of $He(2^{3}S)$ from the upper thermosphere to the exosphere was drastically changed by forcing both from the lower atmosphere and from the space. $He(2^{3}S)$ measurements will definitely improve our understanding of thermosphere-ionosphere coupling system and extend the coverage of space weather forecasting up to the exobase.